

# Automated Processing of Internal Ice Layer Dip in Radio-Echo Soundings

Louise Sime, Richard Hindmarsh, and Hugh Corr

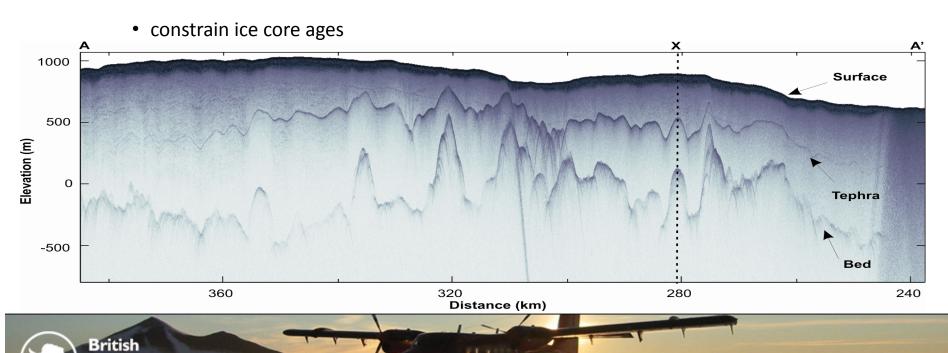




#### **ARESP**

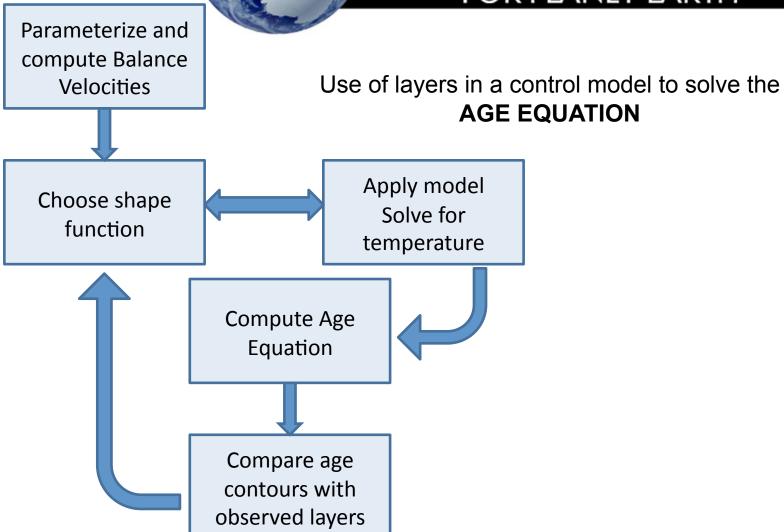
#### Automated RES processing

- Internal layers are considered to be isochronous
- They provide a picture of the age structure of the ice and used to:
  - · investigate changes in ice flow
  - flow modelling studies
  - calculate past accumulation rates



#### **ARESP**



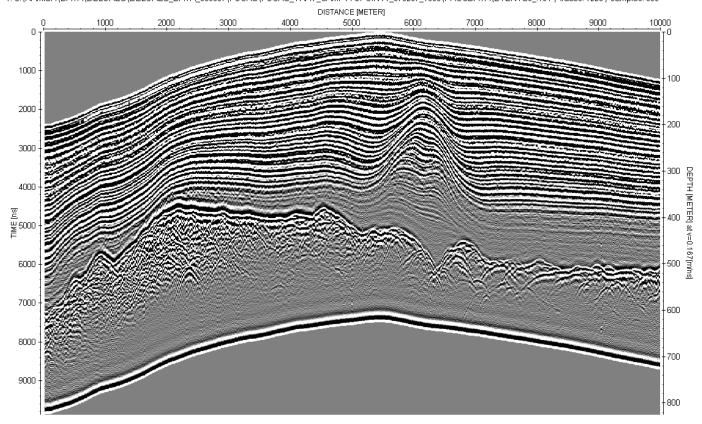






Problem: layers are not 'continuous' (question of length) Picking is labour intensive even with automatic (semi) techniques

1. C:\HAMISH\DATA\DELORES\DELORES\_DATA\_050307\FUCHS\FUCHS\_TRAV\_CAMP1TOPOINT1\_070207\_1335\PROCDATA\EVENTS0\_10T / traces: 1226 / samples: 989

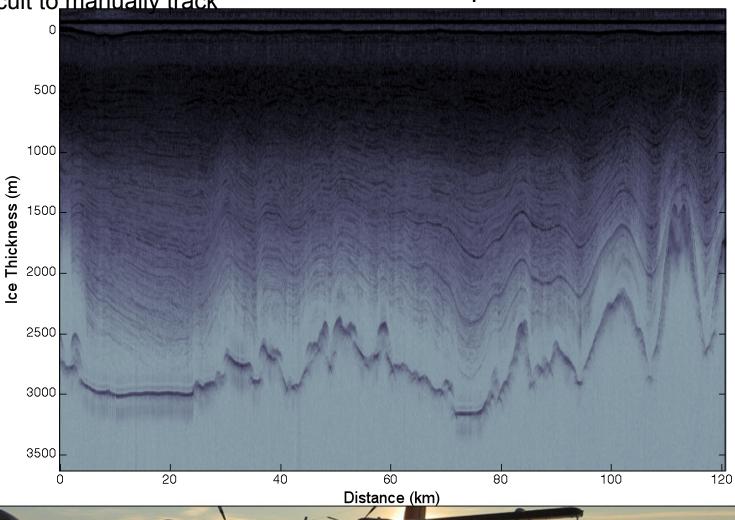




Layers appear continuous

But are difficult to manually track

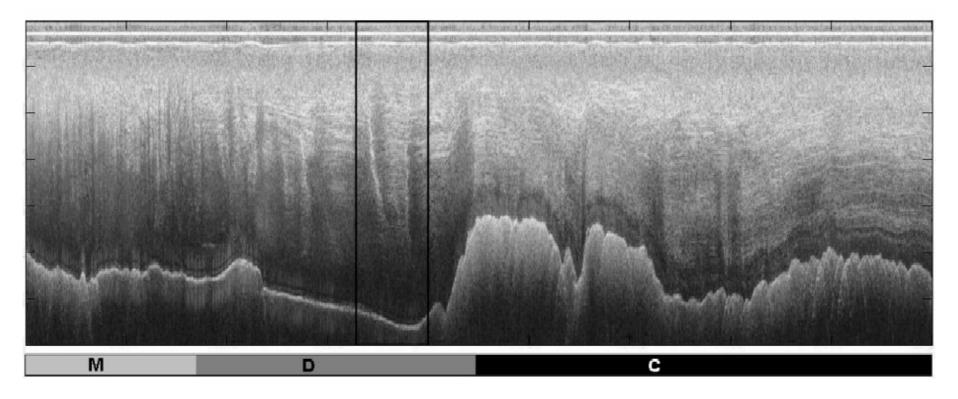
W22 Talos Dome to Europa







#### In regions of fast flow layers are not continuous



45 km section showing M (missing), D (discontinuous) and C (continuous) layering





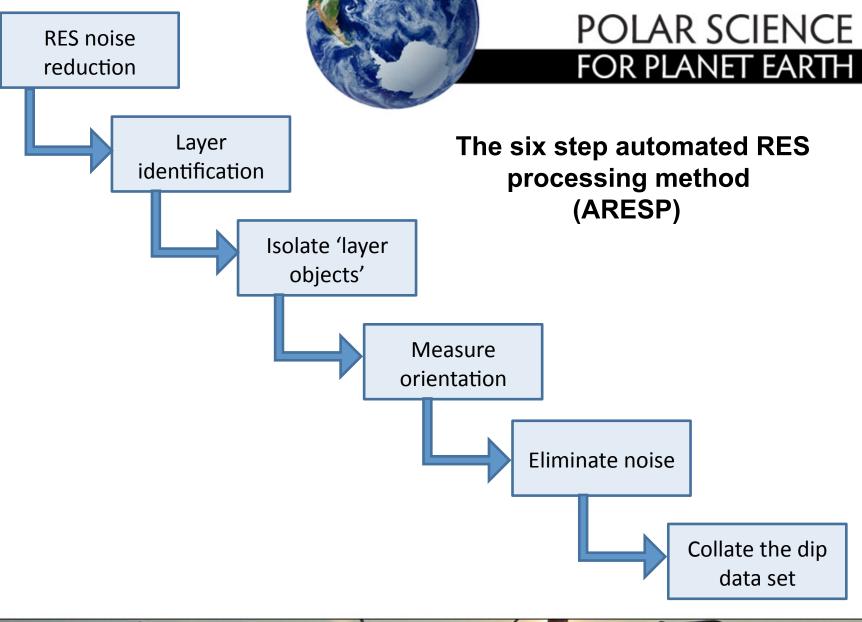
**Solution:** *Parrenin and others (2006)* showed that identical flow information is contained in the apparent dip of the layers.

- Dip datasets can be derived from localised layer information without requiring long-distance layer continuity
- Continuous apparent layer dips can also be integrated to provide synthetic isochrones
- ❖ In most circumstances, individually picked layers or continuous sections of apparent layer dip can equally well be compared with model outputs

Parrenin, F., R. Hindmarsh and F. Remy, 2006. Analytical solutions for the effect of topography, accumulation rate and lateral flow divergence on isochrone layer geometry, J. Glacio., 52(177).

They developed a new coordinate transform moving from physical space to a system of (logQ, log $\omega$ ), in which the streamlines are straight lines. Where Q is the horizontal flux and  $\omega$  is the flux shape function







Step 1: Noise reduction

Horizontal moving-average window of 100 metres

Decimate by a factor of 10 Vertical averaging with length  $\lambda_{layer}$  /2 to give  $P_{avg}$ 1500 Depth (m) 2500 3000 3500 4000 250 230 240 270 Distance (km)





#### **Step 2**: Obtain a Binary Layer Dataset

Directly threshold using a localised threshold

First again vertically average with length  $(\lambda_{layer} \times 2)$  to produce  $P_{avg2}$ 

Then,  $B = sign(P_{avg} - P_{avg2})$ 

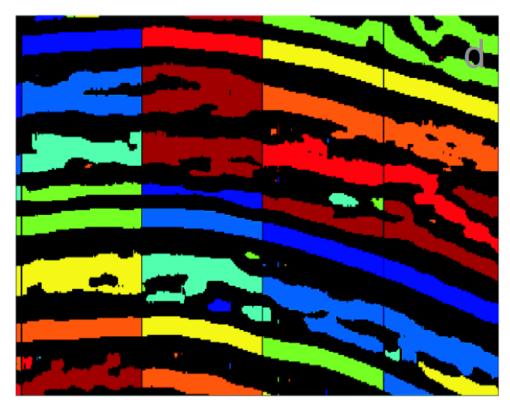




Step 3: Isolate layer objects

Horizontally separate the binary array into thin vertical stripes (~ 100 metres)

An object is adjacent values in a stripe with the same value

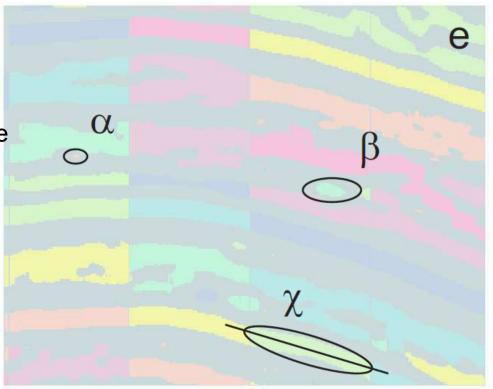






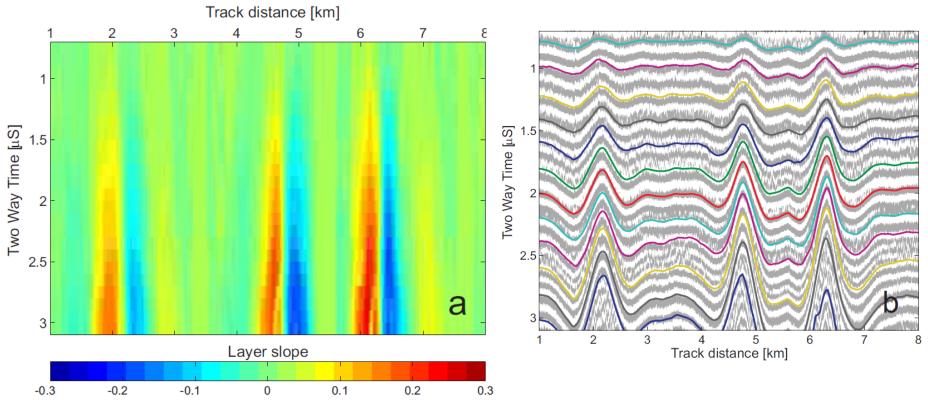
**Step 4**: *Measure 'layer objects'*Position, area, major and minor axis length, and orientation

Step 5: Eliminate invalid 'layer objects'
For example α and β will not give accurate angles



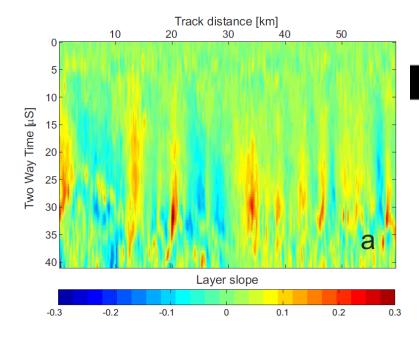


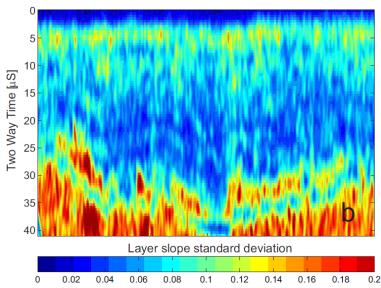




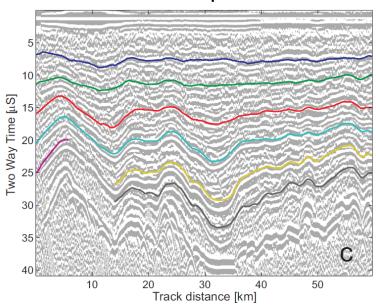
Ground based example. (a) The collated ARESP layer slope data (b) radargram overlaid with synthetic 'layers'







#### Airborne data example



Note: gray image is a standard score = (raw – mean)/std





#### CONCLUSION

- ARESP is a novel, automated and robust method.
- ➤ Continuous layers along the whole length of the RES section are not necessary.
- ➤ Because of its efficiency it will enable regional rather than localized studies.
- Work in progress:
  - On the method to automate variable specification for any RES dataset.
  - For direct model-observational comparison or for inverse modelling, there is an issue of how little observational data is required.

